

**Melt modeling:
Temperatur-index models vs.
Energy balance models**

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(Slides: modified from R. Hock)

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How to model melt ?

- 1. **Physically based energy-balance models:**
each of the relevant energy fluxes at the glacier surface is computed from physically based calculations using direct measurements of the necessary meteorological variables
- 2. **Temperature-index or degree-day models:**
melt is calculated from an empirical formula as a function of air temperature alone

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Melt Modelling

Temperature index | **Energy balance**

Data requirements Model sophistication

Input data:
Air temp | Air temp, humidity, wind speed, radiation, ****

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Mass balance models

0-Dim | elevation bands | fully distributed

1 Temp-index regression | 2 Temp-index or simplified energy balance | 3 Energy balance

Increasing model sophistication

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Temperature-index melt models

- Assume a relationship between air temperature and melt: $M=f(T)$, $M=f(T^*)$

Relationship melt - air temperature

Monthly ablation rate mm water/day | Monthly mean temperature deg | $R=0.93$

Daily ablation mm d⁻¹ | Daily temperature deg

Data by R. Braithwaite
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Temperature-index melt models

- Assume a relationship between air temperature and melt: $M=f(T)$, $M=f(T^*)$

Positive degree-day sum

$PDD = \sum T^+$

Relationship melt - degree-day sum

Ice ablation mm water | Degree-days deg d

Monthly equivalent

Legend: GCM, NBG, IMAU, AWI, EGIS

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Physical basis of temp-index models

Air temperature directly affects several components of the surface energy balance

Longwave incoming rad: $L_{\downarrow} = \epsilon \sigma T^4$

Sensible heat flux: $f(T)$

Latent heat flux: $f(e(T))$

$$Q_M = G(1-\alpha) + L_{\downarrow} - L_{\uparrow} + Q_H + Q_L$$

- Longwave incoming radiation (L_{\downarrow}) is the largest contribution to melt (~70%) (Ohmura, 2001: Physical basis of temperature-index models)
- L_{\downarrow} has low variability compared to other fluxes

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Spatial variability of degree-day factors

Calculation of degree-day factors for various points on the Greenland ice sheet with an atmospheric and snow model (thesis Filip Lefebre)

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Degree-day factors [mm/day/K]

Spatial and diurnal variation
Derived from energy balance modeling

$$M = DDF_{ice/snow} \cdot T^+$$

Hock, 1999, J. Glaciol.

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Modified temperature-index model

Including potential direct solar radiation

Classical degree-day factor: $M = DDF_{ice/snow} \cdot T^+$

Including pot. direct radiation: $M = (MF + a_{ice/snow} \cdot DIR) \cdot T^+$

Model introduces:

- a spatial variation in melt factors
- a diurnal variation in melt factors

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Modified temperature-index model

Simulated cumulative

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Distributed temp-index model

$$M = TF \cdot T + AF \cdot (1-\alpha) \cdot I \quad T > T_T$$

Temperature: Computed as function of accumulated maximum daily temperature since snowfall (Brock et al. (2000))

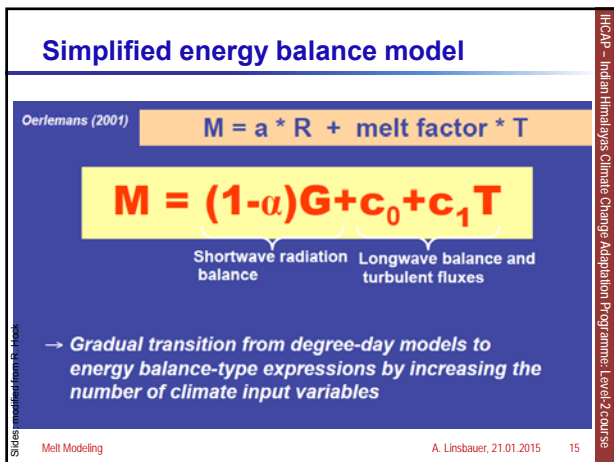
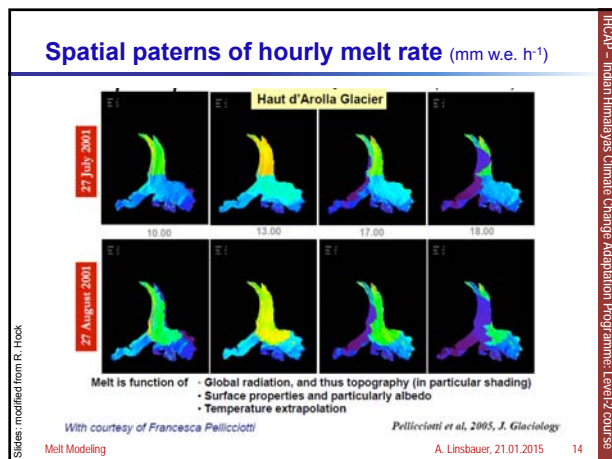
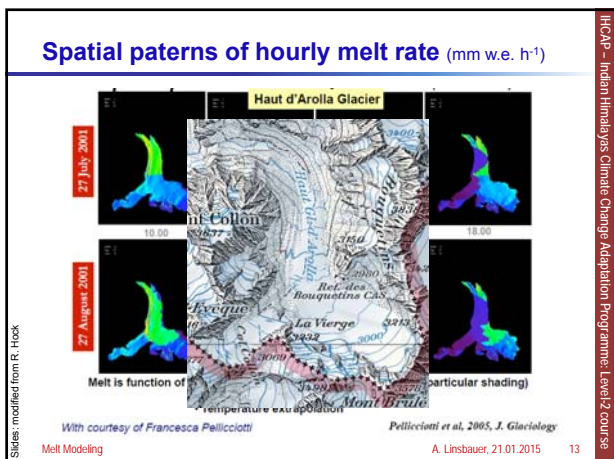
Albedo: Computed as function of daily temperature range

Incoming shortwave radiation: Computed as function of daily temperature range

Model only requires air temperature
Global radiation and albedo parameterized

Pellicciotti et al, 2005, J. Glaciology

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Temperature-index versus energy balance

	Temperature index	Energy balance
Advantages	<ul style="list-style-type: none"> Wide availability of Temp-data Easy interpolation and forecasting Good model performance Computational simplicity 	<ul style="list-style-type: none"> Physical based – describe physical processes more adequately Projections more reliable
Shortcomings	<ul style="list-style-type: none"> Empirical, not physically based DDF vary, works on 'average conditions' Does not work in tropics Model parameter stability under different climate conditions? 	<ul style="list-style-type: none"> Large data requirements (often not available)

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